



TEMPORAL AND SPATIAL VARIATIONS OF DISCRETE MAGNETOSPHERIC OSCILLATIONS DETECTED BELOW 1mHz

**Jannie Marfaing^a, Elisabeth Pozzo di Borgo^b, Matthew Yedlin^c,
Jean Jacques Bois^d, Jonathan Fraser^c, Georges Waysand^d,
Stephane Gaffet^{d,e}, Rémi Blancon^b and Alain Cavailleou^d**

^a Universités d'Aix-Marseille, IM2NP, case 142, Faculté des Sciences de Saint Jérôme,
F13397 Marseille Cedex 20 (jannie.marfaing@im2np.fr)

^b Université d'Avignon et des Pays de Vaucluse, UMR 1114 INRA UAPV, rue Louis
Pasteur F-84000 Avignon

^c University of British Columbia, Dept. of Electrical and Computer Engineering, 2356
Main Mall, Vancouver, BC, V6T 1Z4, Canada

^d LSBB, Observatoire de la Côte d'Azur, UNS/CNRS/OCA, La Grande Combe, F-
84400 Rustrel

^e GEOAZUR, UNS/CNRS/OCA, 250 rue Albert Einstein, Sophia-Antipolis, F-06560
Valbonne

Keywords: *Earth's eigenmodes – ionosphere-Earth coupling –
magnetic detection – interferences – SQUID magnetometer*

• **Low frequency geomagnetic pulsations are detected below 2 mHz, using the underground ultra low-noise 3D superconducting magnetometer system [SQUID]² (SQUID with Shielding Qualified for Ionosphere Detection) in the Laboratoire Souterrain Bas Bruit in Rustrel (LSBB - France).** The detection of the Earth's eigen modes in the magnetic quiet background of the ionosphere has been established [1] **BUT** From experimental data collected from ground-based stations and from sky satellites devoted to the study/survey of the magnetosphere, several power spectra of magnetic pulsations, entering the magnetosphere, have been identified.

• **Scientific questions:**

- **Nature of these waves: continuous? irregular? polarized? ...**
- **Origin of these waves**
- **Evaluation of the energy transfert**
- **autoconsistant theoretical model**

• **Need:**

- **Identification of the discrete frequencies spectrum**
- **Characterization**

- **Analysis of the 3D-magnetograms made for 72 consecutive hours in absence of major earthquakes and magnetic storms: same conditions as for detection of Earth's eigenmodes**
- **Use of 3 techniques:**
- **the Fast Fourier Transform (FFT) and Discrete Fourier Transform (DFT) : frequency signature**
- **DFT allows an adjustable resolution for a set of selected frequency peaks using the 2nd order Goertzel's algorithm (Beraldin and Steenaart, 1989)**
- **the Stockwell Transform (ST), extension of the short time Fourier Transform and a variant of the continuous wavelet transform: time-frequency décomposition.**
- **Advantage: modern tool for non-periodic and/or non-stationary signals with evidence of the discontinuous character of the signal.**

Discrete Fourier Transform (DFT): FT + 2nd order Goertzel algorithm
Resolution: $df = 3.86 \mu\text{Hz}$

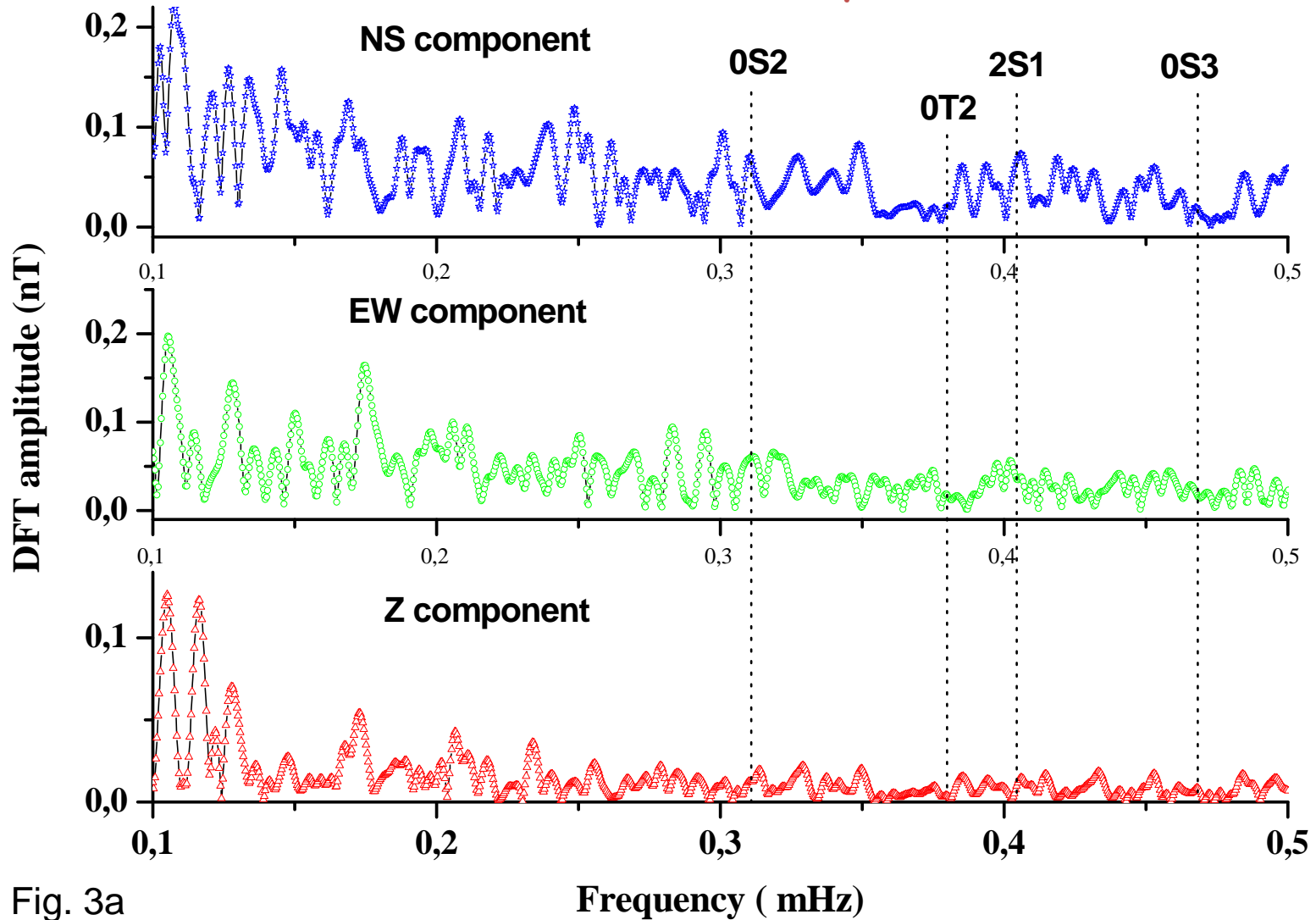


Fig. 3a

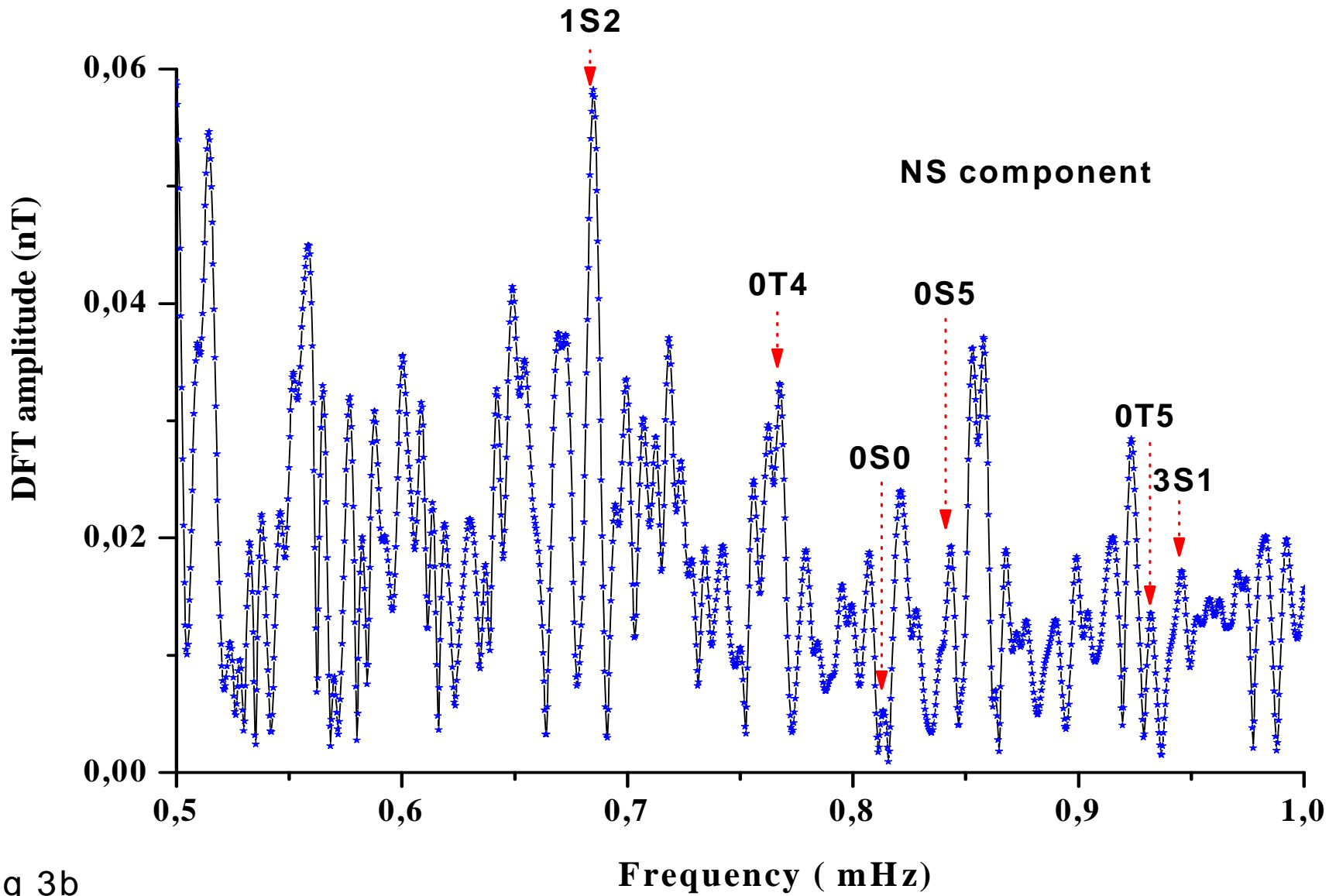
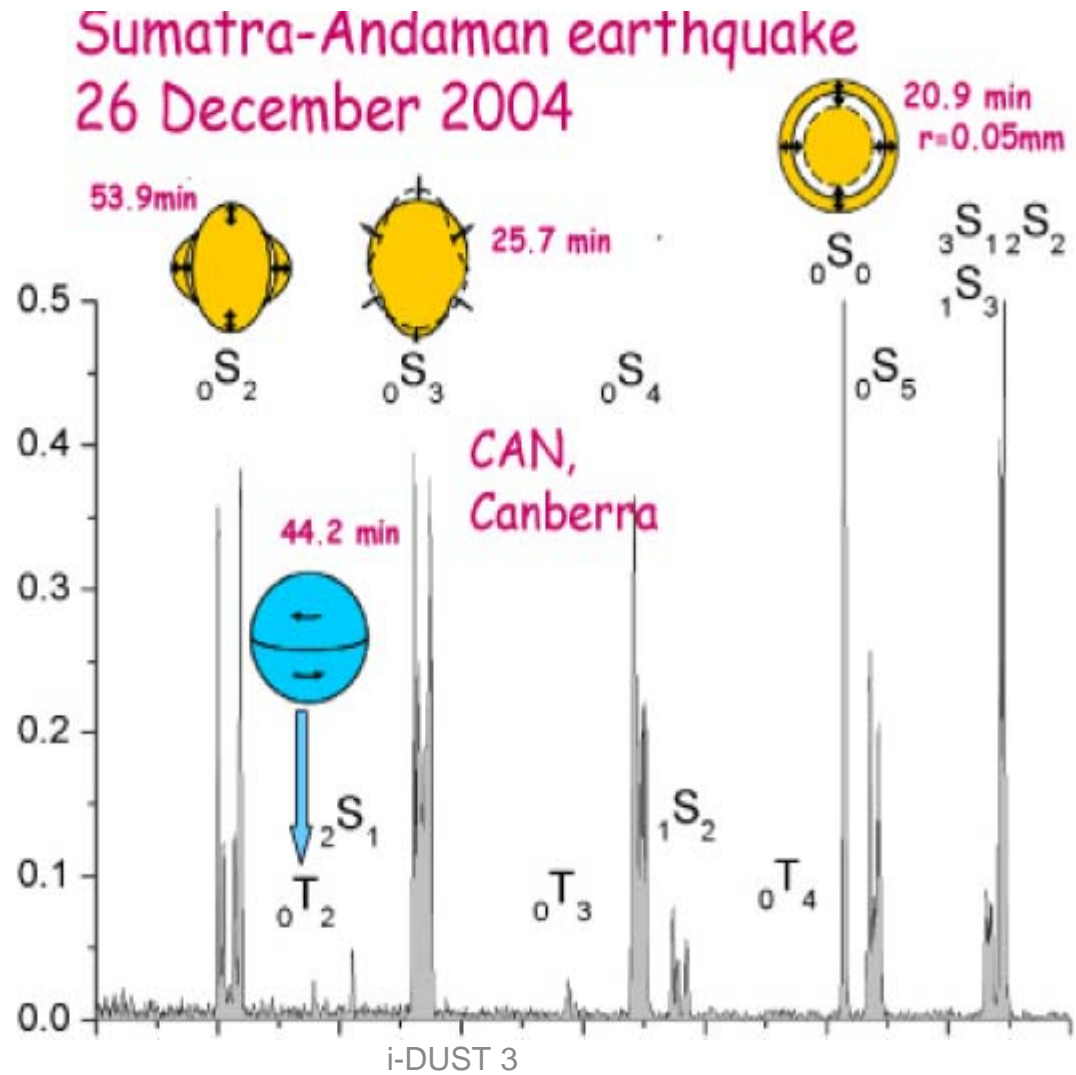


Fig 3b

Comparison with the seismic detection during a major earthquake ($M_w = 8$)



Stockwell Transform (ST): is a variant of the continuous wavelet transform given by

$$W(b, a) = \int_{-\infty}^{\infty} f(x) \frac{1}{\sqrt{|a|}} \Phi\left(\frac{x-b}{a}\right)^* dx \quad (1)$$

The Stockwell transform uses a particular complex mother wavelet – hence the conjugate.

The Stockwell transform defined in Stockwell et al. 1996, is given by

$$S(b, \xi) = \int_{-\infty}^{\infty} f(x) |\xi| \frac{e^{-\frac{(x-b)^2 \xi^2}{2}}}{\sqrt{2\pi}} e^{-2\pi i \xi x} dx$$

where b corresponds to time, t and ξ corresponds to the frequency. If we factor out the phase modulation corresponding to

$$S\left(b, \frac{1}{a}\right) = \frac{1}{\sqrt{|a|}} e^{-2\pi i \frac{1}{a} b} \int_{-\infty}^{\infty} f(x) \frac{1}{\sqrt{|a|}} \frac{e^{-\frac{(x-b)^2}{2a^2}}}{\sqrt{2\pi}} e^{-2\pi i \frac{1}{a}(x-b)} dx$$

$$= \frac{1}{\sqrt{|a|}} e^{-2\pi i \frac{1}{a} b} \int_{-\infty}^{\infty} f(x) \frac{1}{\sqrt{|a|}} \left[\frac{e^{-\frac{1}{2}\left(\frac{x-b}{a}\right)^2}}{\sqrt{2\pi}} e^{-2\pi i \left(\frac{x-b}{a}\right)} \right] dx$$

where the term in the square brackets is of the form $\frac{1}{\sqrt{|a|}} \Phi\left(\frac{x-b}{a}\right)$ *

when F is chosen to be the complex Morlet wavelet given by:

$$\Phi(x) = \frac{1}{\sqrt{2\pi}} e^{-\frac{1}{2}x^2} e^{2\pi i x}$$

Thus, we see that **the Stockwell transform is a continuous wavelet transform given by**

$$S\left(b, \frac{1}{a}\right) = \sqrt{\frac{1}{|a|}} e^{-2\pi i \frac{b}{a}} W(b, a)$$

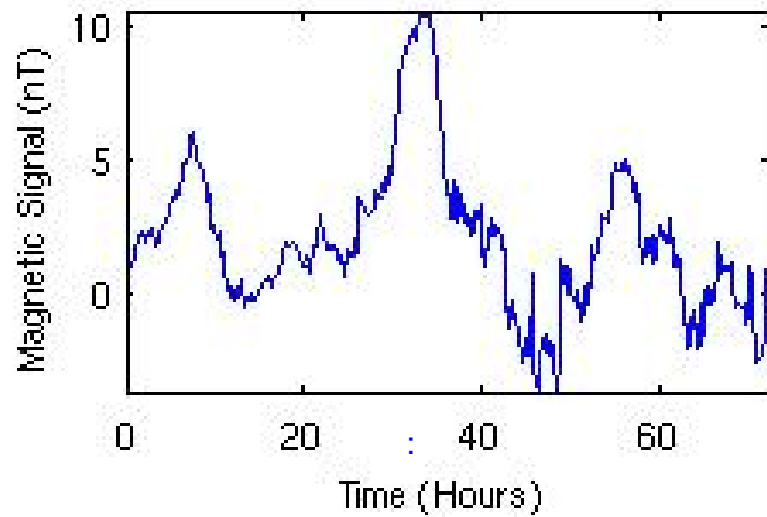
Appropriate low-pass filtering, .

Resampling of the data at 1 minute intervals, which corresponds to a Nyquist rate of 1/120 Hz which is 8.33 mHz.

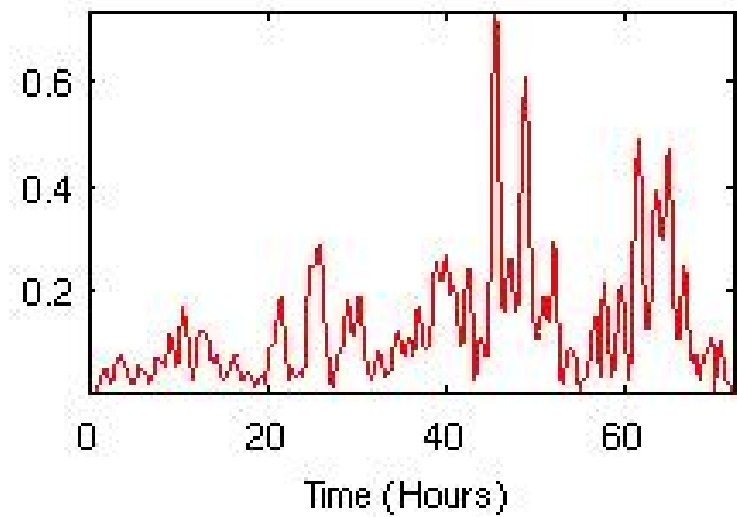
Frequency analysis using the Stockwell transform : between 0.2 and 1 mHz, well below the Nyquist rate for the resampled data.

For the data length of 72 hours, the time series will be over 4000 points long

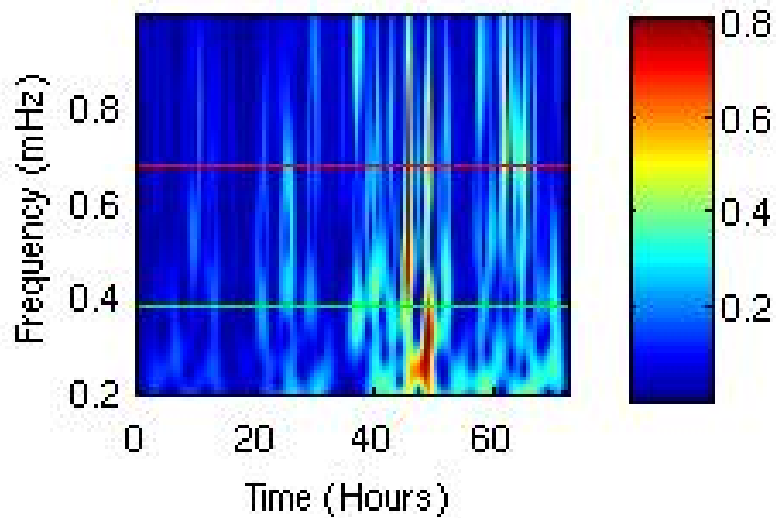
a) NS Time Series



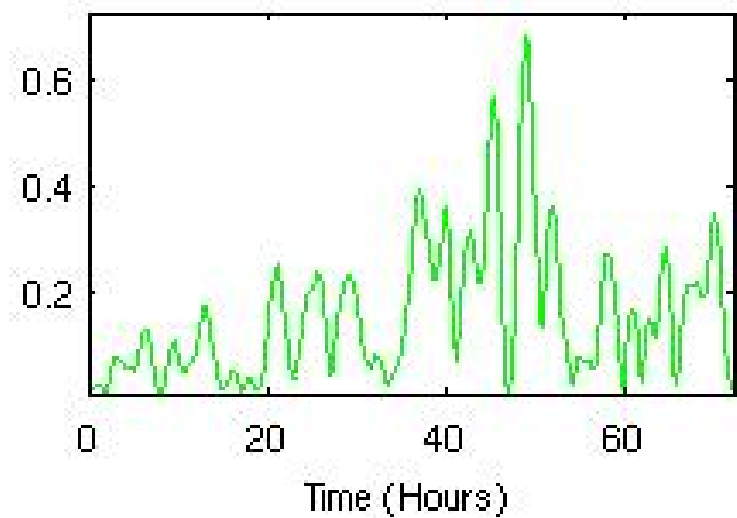
c) 152 mode Stockwell Amplitude

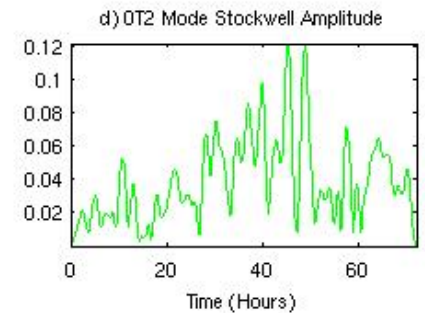
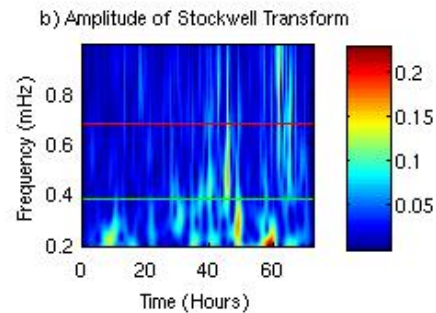
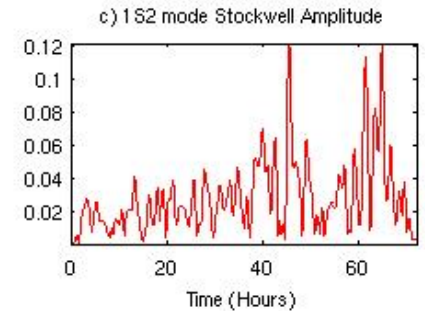
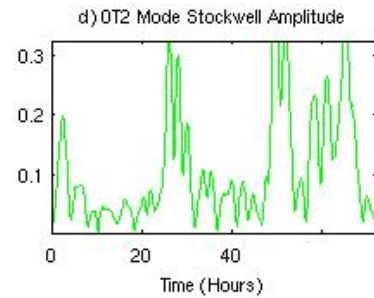
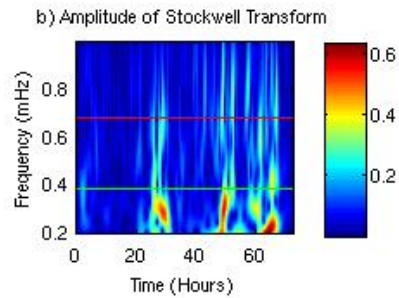
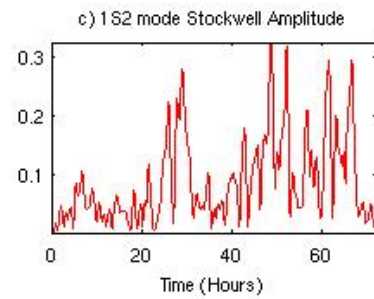
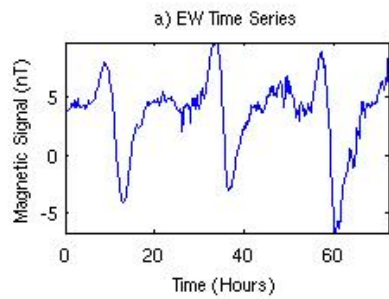


b) Amplitude of Stockwell Transform



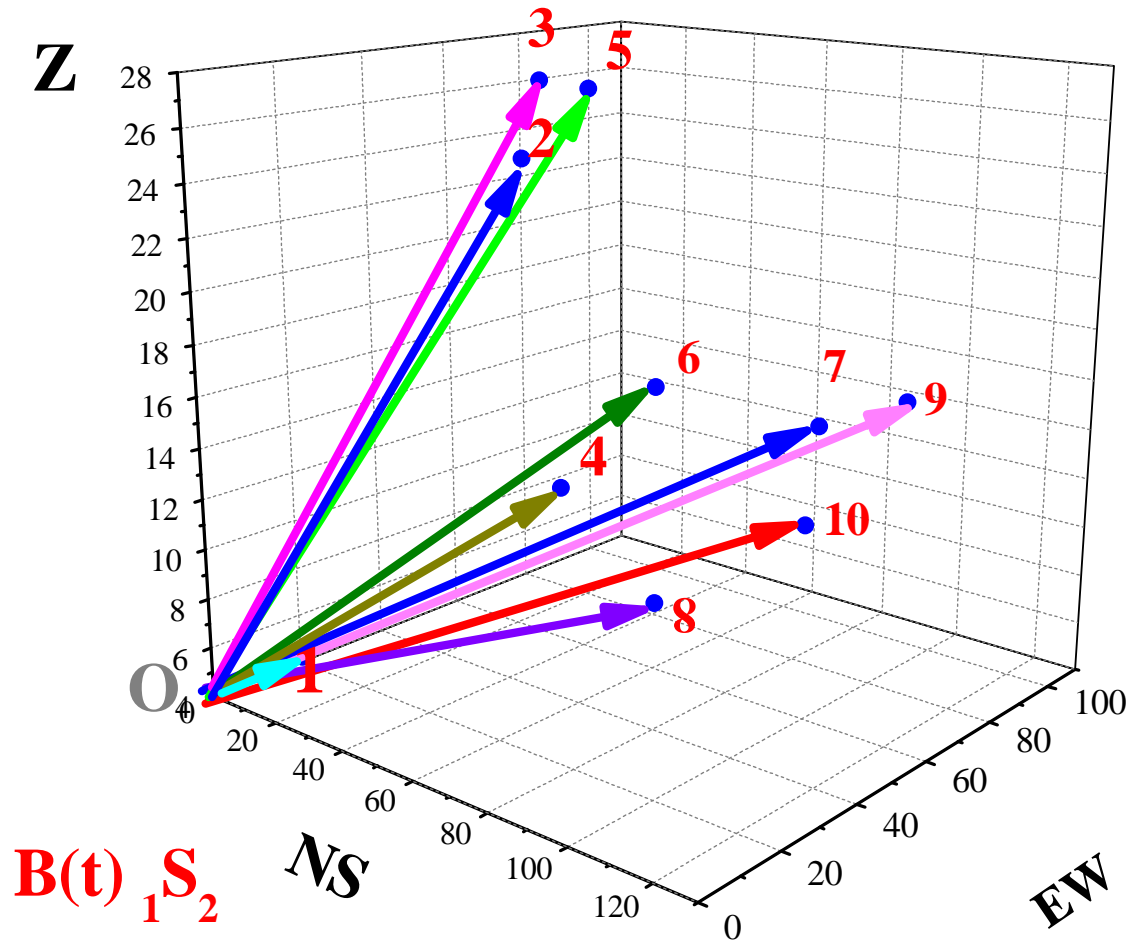
d) 0T2 Mode Stockwell Amplitude





Z

Dependence in time of the mode 1S2 over 10 hours



UNIDENTIFIED PULSATIONS BELOW 1 mHz

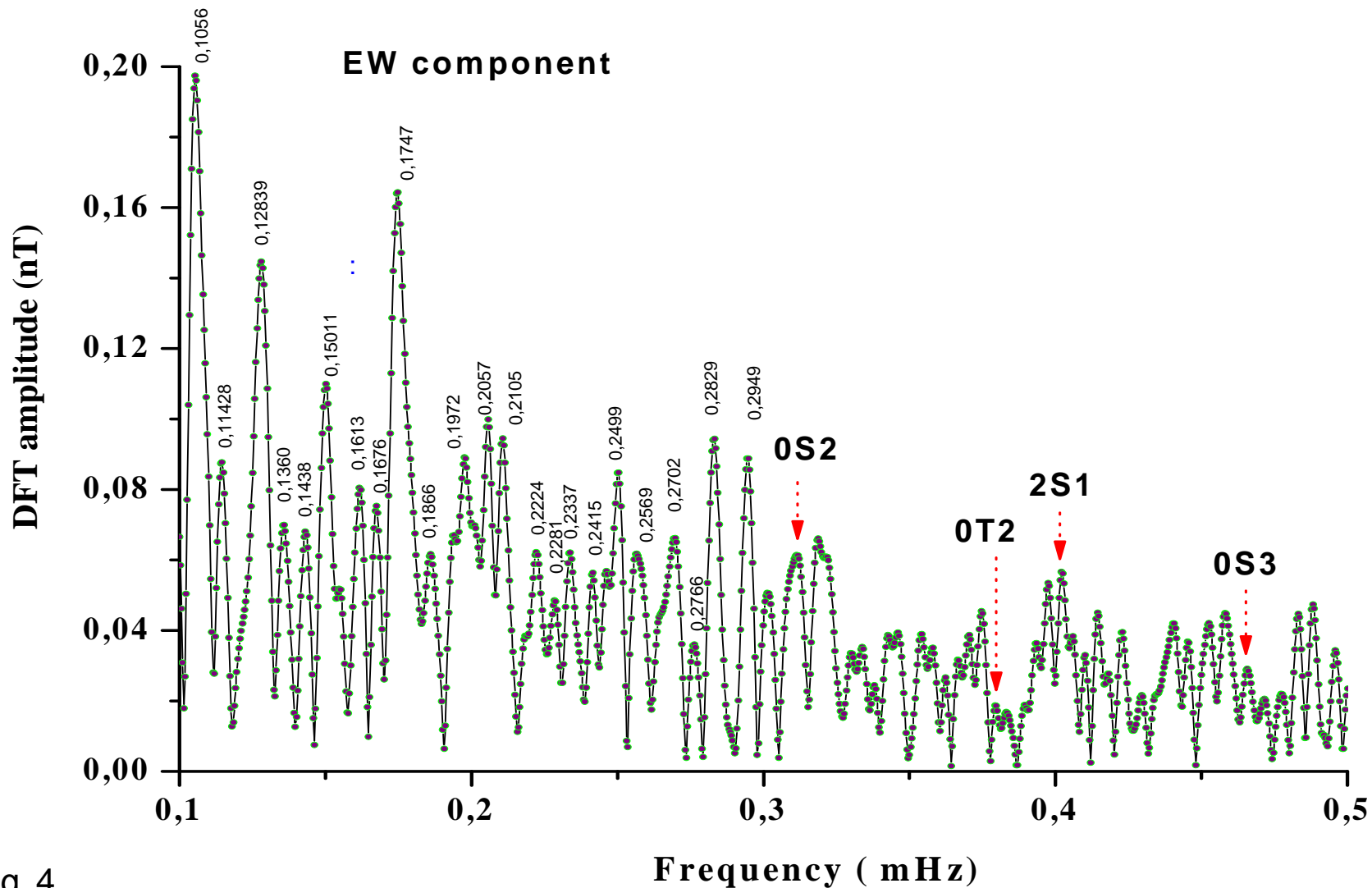


Fig 4

AND ALSO BELOW 0.1 mHz

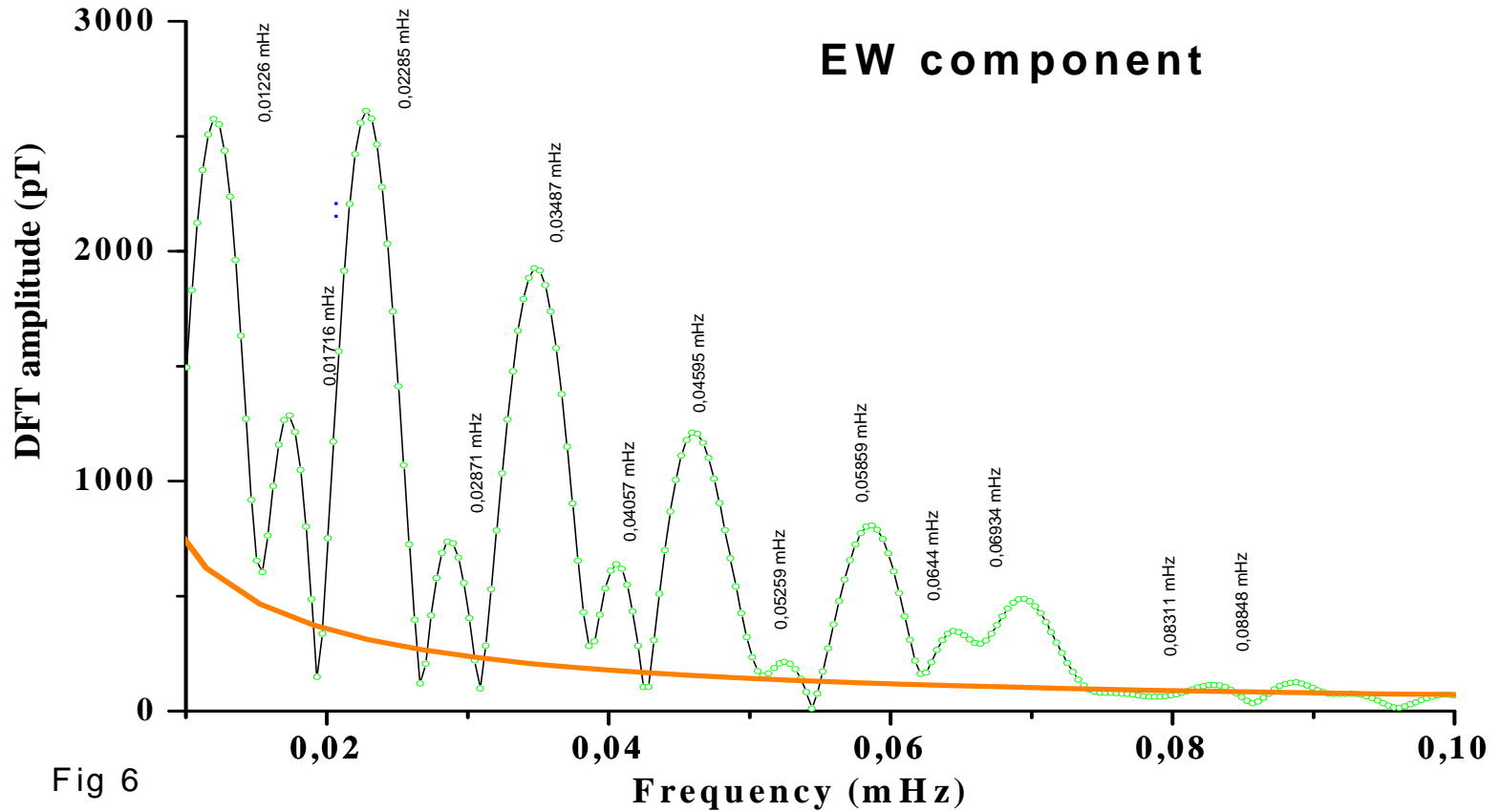
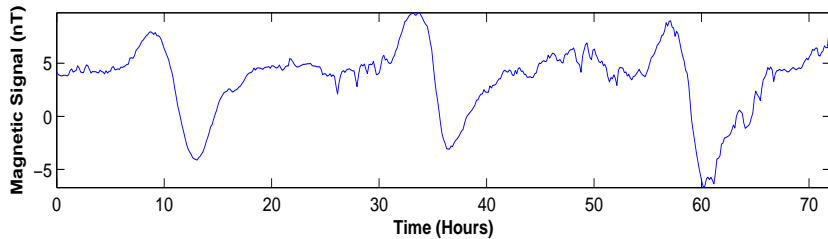
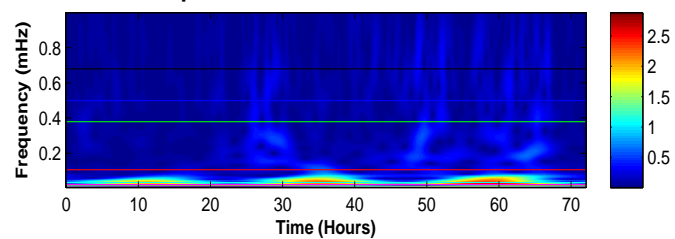


Fig 6

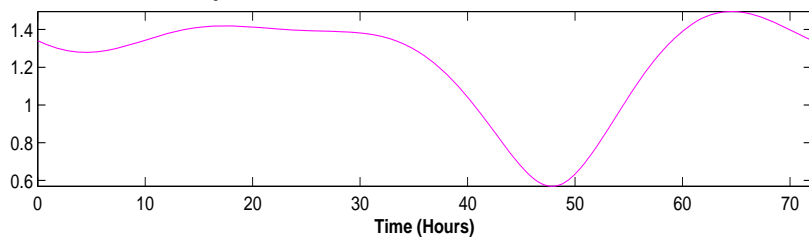
EW Time Series



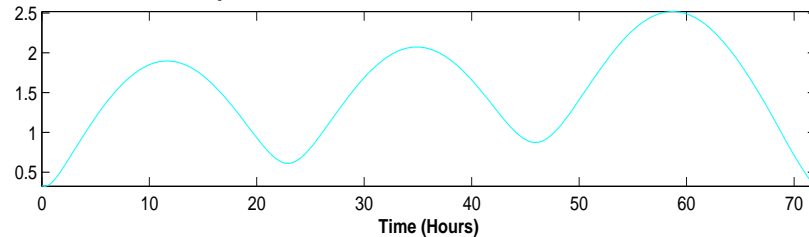
Amplitude of Stockwell Transform



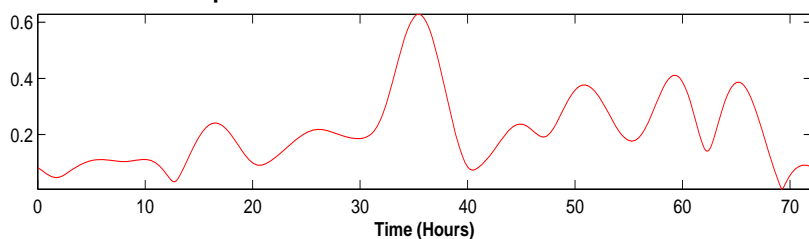
Amplitude of the Stockwell Trace at 0.023 mHz



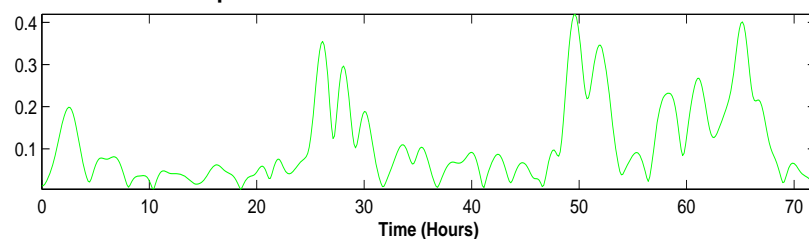
Amplitude of the Stockwell Trace at 0.035 mHz



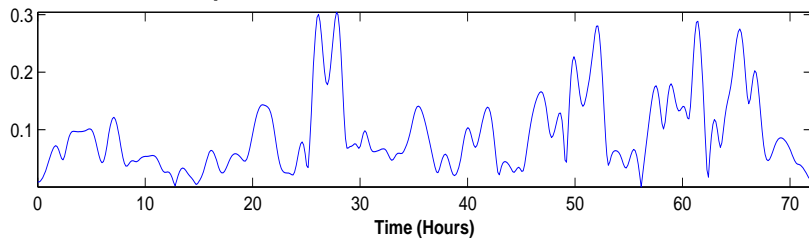
Amplitude of the Stockwell Trace at 0.107 mHz



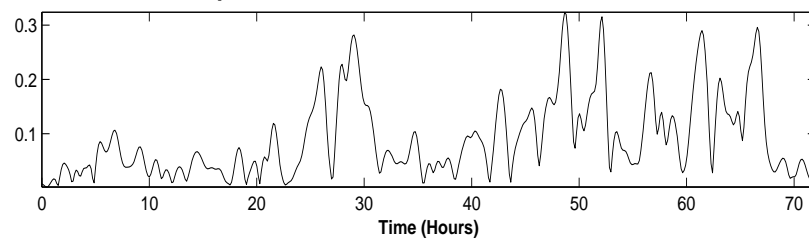
Amplitude of the Stockwell Trace at 0.379 mHz

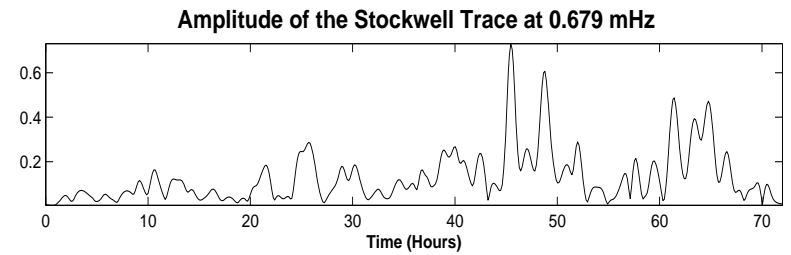
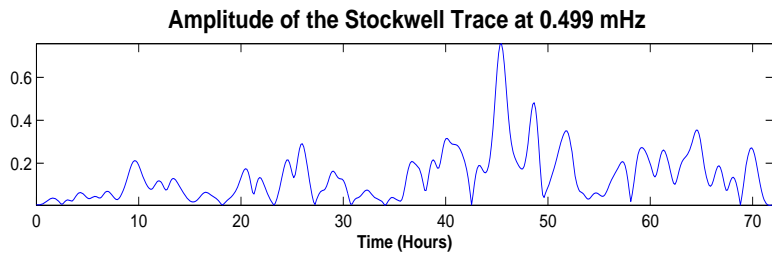
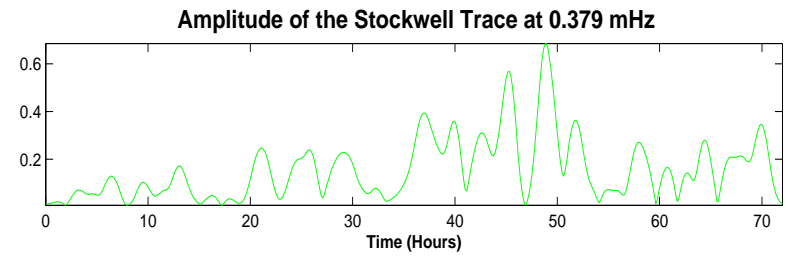
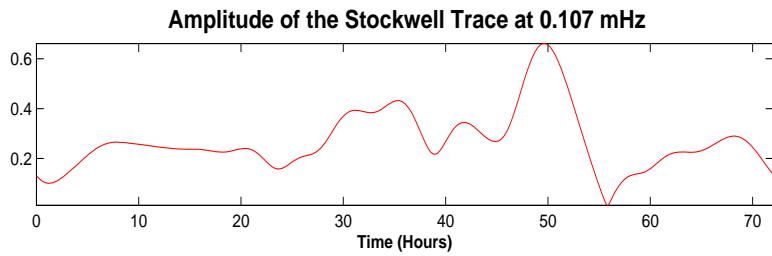
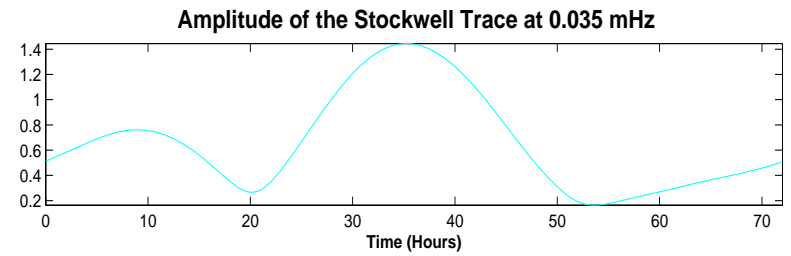
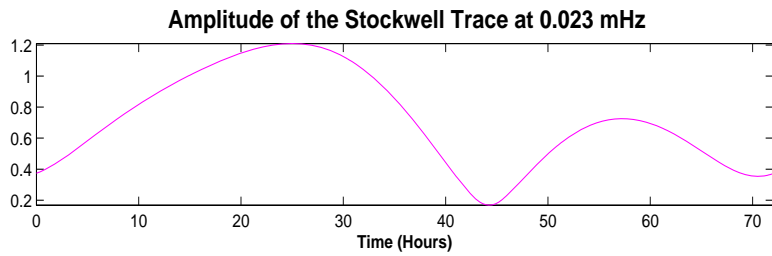
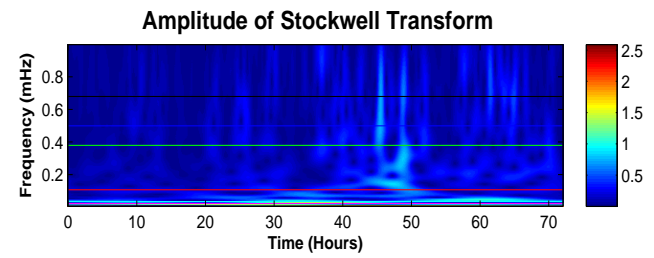
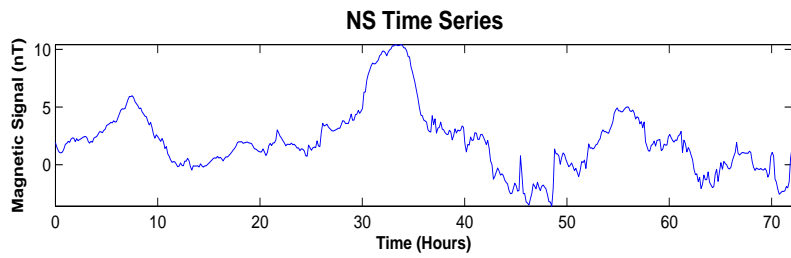


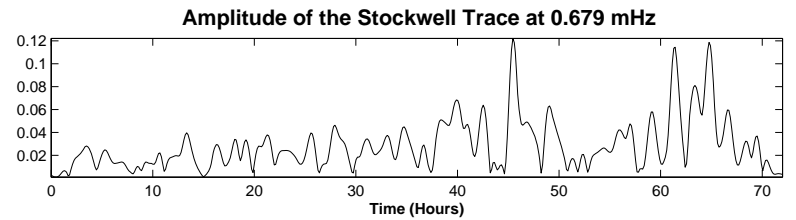
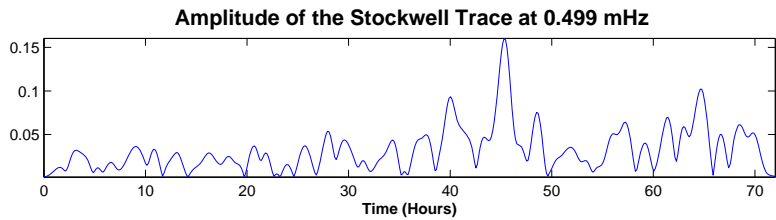
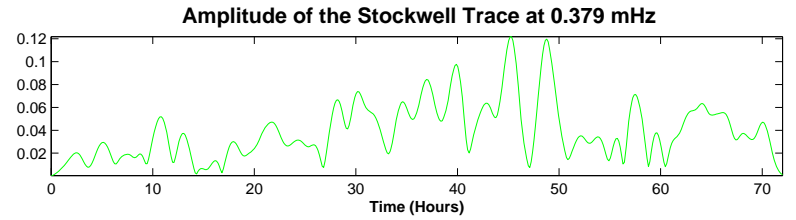
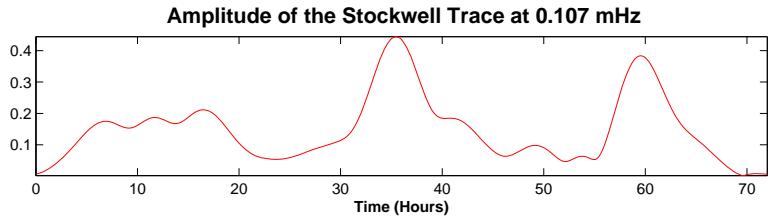
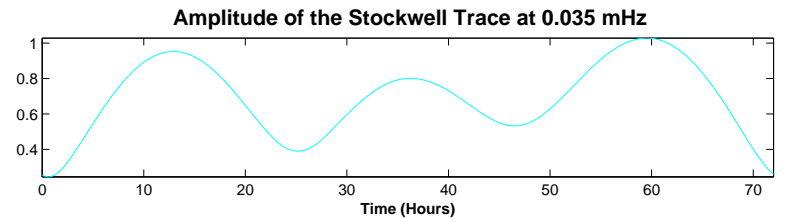
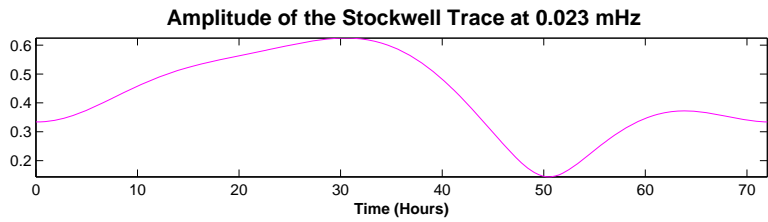
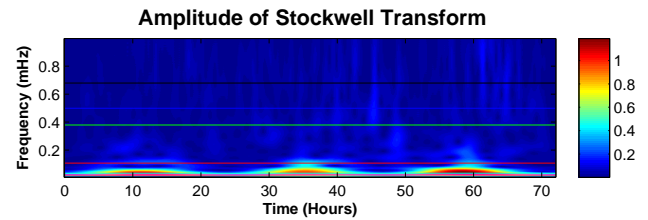
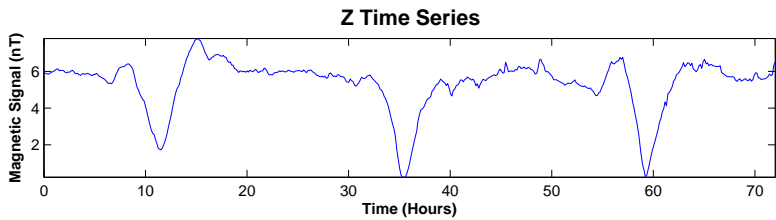
Amplitude of the Stockwell Trace at 0.499 mHz



Amplitude of the Stockwell Trace at 0.679 mHz

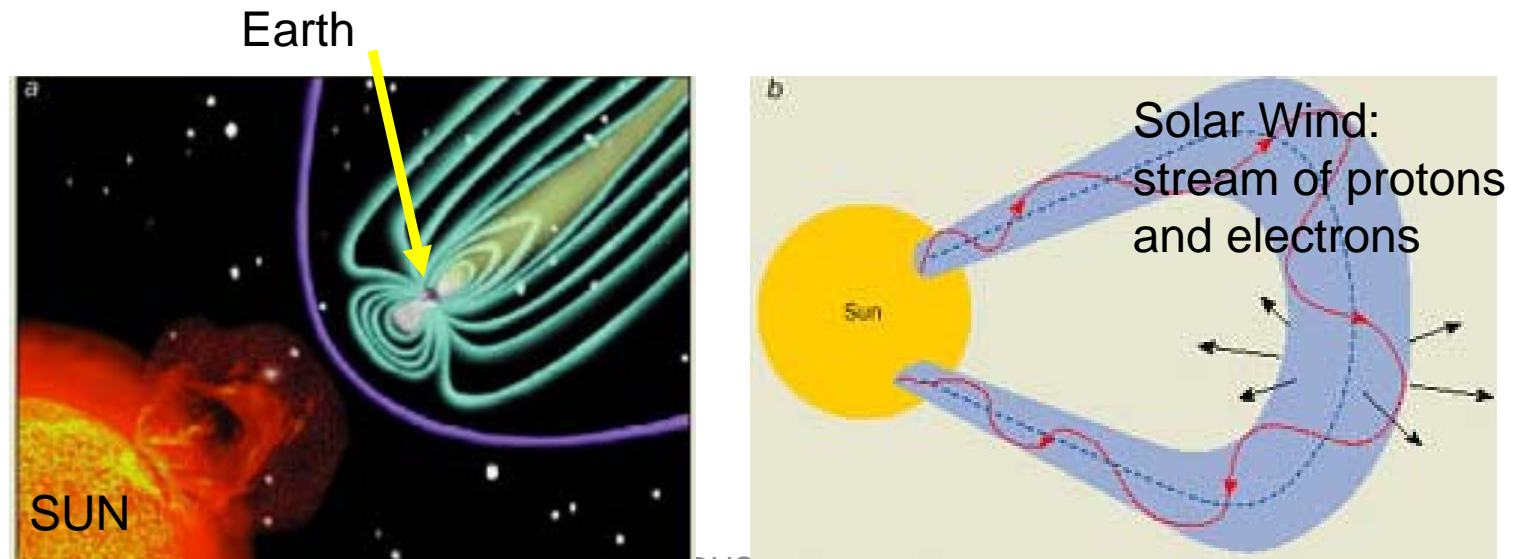






• PROBABLE ORIGIN

- Several sources of ULF waves are known in the millihertz band, originating
- either from the **incessant seismic activity of the Earth** which induces the free oscillation modes of the Earth generally detected by the seismographs between 2 and 7 mHz (Tanimoto 2005)
- or from the **ionosphere and the magnetosphere**, resulting for example from the Kelvin-Helmholtz, Rayleigh-Taylor ...instabilities at the magnetosphere boundary (Kivelson & Russell 1995).



ULF waves originating from ionosphere, especially of type Pc5 are **detected during intense magnetic storms**, due to changes in the solar activity and correlated with solar wind parameters (Vichare *et al.* 2009; Kleimenova & Kozyreva, 2007).

Spectral peaks in the 0.5 – 4 mHz frequency range (Stephenson & Walker 2002; Kepko & Spence 2003): near **0.5-0.6, 0.8, 1-1.5, 1.2-1.4, 1.6, 1.8, 2.0, 2.6,...3.4 mHz.**

However, up to now, no accurate values or complete list of such ULF waves are known from experimental observations or ionospheric models.

Vichare, G., Alex, S. and Lakhina, G.S., *J. Atmospheric and Solar-Terrestrial Physics*, **71**, 2009, 1814-1823, doi:10.1016/j.jastp.2009.06.015

Kleimenova, N.G. and Kozyreva, O. V., “*Physics of Auroral Phenomena*”, Proc. XXX Annual Seminar Apatity, 2007, pp. 30-33.

Stephenson, J.A.E., and Walker, A.D.M., *Geophys. Res. Lett.* **29**, 2002, 121297, doi:10.1029/2001GL014291,2002

Kepko, L. and Spence, H.E., *J. geophys. Res.*, **108** (A6), 2003, 1257; doi : 10.1029/2002JA009676,2003

Another origin can be speculated:

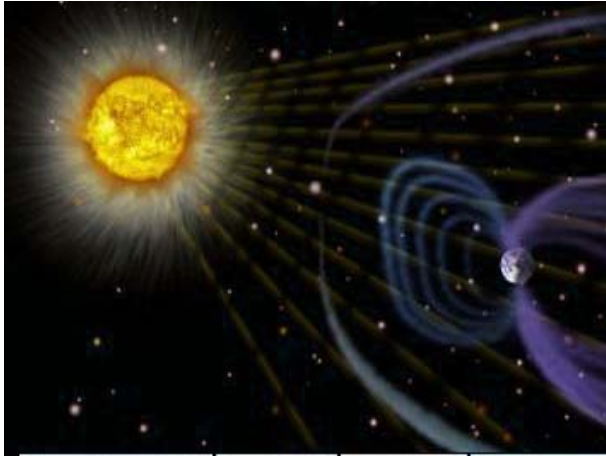
- **beat frequencies** : simple difference between two close Earth's eignemodes frequencies that can be easily identified from the calculated PREM values.
- **50 frequencies, deviating by less than 1% from the eigenmodes of Earth, have been magnetically detected below 2 mHz (54 theoretical values).**
Implicitly, interference exists if the waves present similar amplitude and frequencies.

In the lowest frequency range, from 0.01 to 0.3 mHz, it has been verified that:

- i) **all the unidentified peaks in the range 0.01 to 0.3 mHz may correspond to interferences between [SQUID]²-identified modes, whether S or T modes**

there is an excellent agreement between the observed and “calculated” interferences.

their amplitudes can be high as several modes can present the same difference in frequency, and hence produce constructive signals.



'Earth's magnetosphere is a very large, complex and variable system. This makes the understanding of ULF waves, together with the mechanisms for the energy transfer from space to ground, a very difficult matter,' says Philippe Escoubet, ESA's Cluster and Double Star Project Scientist.

Using the data from GEOS-2 satellite, YOMOTO et al. 1984 have proposed that the Pc3-4 wave energy is convected through the magnetosheath to the magnetopause, transmitted deep into the magnetosphere without significant changes in spectra, and then couple with various hydromagnetic wave modes in the magnetosphere (Ansari 2006).

Other track: Study of integrated flux variations of four lines (N _ O _ Ne) - search of periodicities in the fluxes of the lines (power spectra from FFT and wavelets methods). **Frequencies detected at 0.09 mHz, 0.25 mHz, 0.38 mHz; 0.83 mHz, 1.44 mHz, (N, and O ions) ; 0.17 mHz, 0.92 mHz (Ne ions) [Popescu et al., 2004].**

